

The translational modes of the inner core : why they might not exist

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Strasbourg - 19th november 2014

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Jump conditions at phase changes

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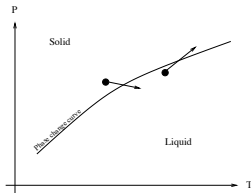
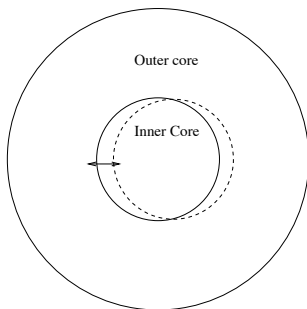
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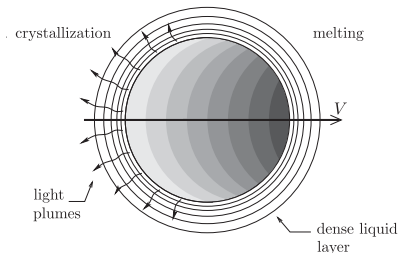
- I. In which situation does the modified jump condition appears
- II. Highlight the problem with a very simple flow-model
- III. A method to investigate boundary conditions

I. In which situation does the modified jump condition appears

1. The Slichter mode : an oscillatory mode



2. The permanent convective translation within the inner core

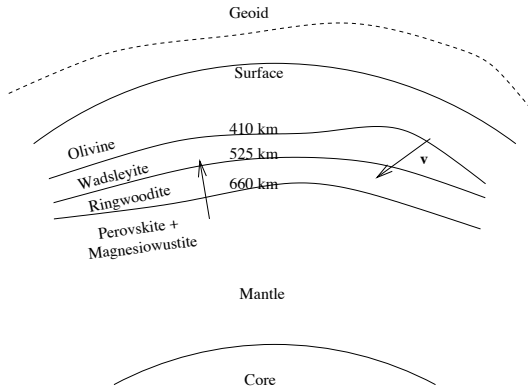


Boundary condition ?

Figure 1. A schematic representation of the translation mode of the inner core, with the grey shading showing the potential temperature distribution (or equivalently the density perturbation) in a cross-section including the translation direction (adapted from Alboussière *et al.* 2010).

Figure: Deguen, 2013

3. The mantle convection



Boundary condition ?

Usual boundary condition :

$$[[\sigma \mathbf{n}]] = 0$$

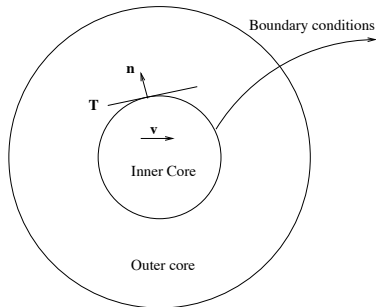
$$\text{Notation : } [[\sigma \mathbf{n}]] := \sigma^+ \mathbf{n} - \sigma^- \mathbf{n}$$

Spherical geometry :

$$\sigma_{rr}^+ = \sigma_{rr}^- \quad \sigma_{r\theta}^+ = \sigma_{r\theta}^- \quad \sigma_{r\phi}^+ = \sigma_{r\phi}^-.$$

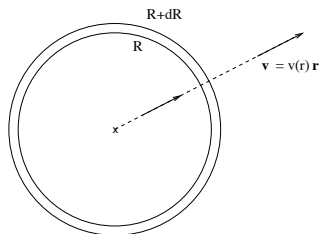
I will show that, at a permeable interface, like a phase change :

$$[[\sigma \mathbf{n}]] = \frac{2\gamma}{R} \mathbf{n} - \nabla_T \gamma.$$



II. Highlight the problem :
a very simple flow-model (spherical)

A simple model to highlight the problem

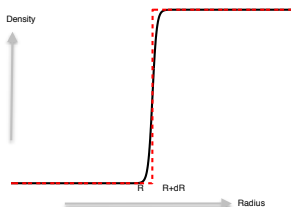


Stokes Equation, viscous linear quasi-static fluid :

$$\nabla \cdot (\rho \mathbf{v}) = 0,$$

$$\nabla \cdot \boldsymbol{\sigma} = 0,$$

$$\boldsymbol{\sigma} = -p\mathbf{I} + \eta \left(\nabla \mathbf{v} + (\nabla \mathbf{v})^T \right) + \lambda (\nabla \cdot \mathbf{v})\mathbf{I}.$$



Radial symmetry :

$$\mathbf{v} = v(r)\mathbf{r},$$

$$\rho = \rho(r).$$

Not essential but simplifies the proof : $\eta = \text{constant}$.

A simple model to highlight the problem

$$\nabla \cdot (\rho \mathbf{v}) = 0,$$

$$\nabla \cdot \boldsymbol{\sigma} = 0,$$

$$\boldsymbol{\sigma} = -p\mathbf{I} + \eta \left(\nabla \mathbf{v} + (\nabla \mathbf{v})^T \right) + \lambda (\nabla \cdot \mathbf{v}) \mathbf{I}.$$

$$\partial_r(\rho v r^2) = 0$$

$$\partial_r \sigma_{rr} + \frac{1}{r} (2\sigma_{rr} - \sigma_{\theta\theta} - \sigma_{\phi\phi}) = 0$$

$$\sigma_{rr} = -p + 2\eta \partial_r v + \lambda \partial_r(r^2 v)/r^2$$

$$\sigma_{\theta\theta} = \sigma_{\phi\phi} = -p + 2\eta v/r + \lambda \partial_r(r^2 v)/r^2$$

$$\partial_r \sigma_{rr} + 4\eta \partial_r \left(\frac{v}{r} \right) = 0.$$

Solution :

$$v = \frac{\text{Cst}}{\rho r^2}$$

$$\sigma_{rr} + 4\eta \frac{v}{r} = \text{Cst}'.$$

If $\rho \rightarrow$ discontinuous :

$$[[\rho v]] = 0$$

$$[[\sigma_{rr}]] = -\frac{4\eta}{R} [[v]]$$

Young-Laplace law :

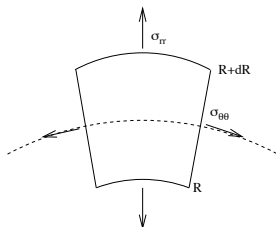
$$[[\sigma_{rr}]] = \frac{2\gamma}{R}.$$

but with

$$\gamma = -2\eta [[v]] = -2\eta \rho v [[1/\rho]].$$

The surface tension γ depends on the flow.

We try to understand why
Think of a thin balloon or a soap bubble



$$\sigma_{rr} = -p + 2\eta \partial_r v + \lambda \partial_r (r^2 v) / r^2$$

$$\sigma_{\theta\theta} = \sigma_{\phi\phi} = -p + 2\eta v / r + \lambda \partial_r (r^2 v) / r^2$$

Tectonics, deviatoric stress : push in
one direction $\leftrightarrow$ pull perpendicular

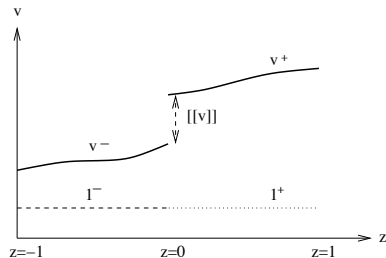
$$\sigma_{\theta\theta} = \sigma_{rr} - 2\eta (\partial_r v_r - v_r / r).$$

$$\implies \sigma_{\theta\theta} \rightarrow \infty.$$

on the boundary.

Surface tension !

III. A method to investigate boundary conditions : discontinuous functions (distributions)



Velocity written in a discontinuous form :

$$\mathbf{v} = \mathbf{v}^- \mathbb{1}^- + \mathbf{v}^+ \mathbb{1}^+,$$

i.e.

$$[[\mathbf{v}]] = \mathbf{v}^+ - \mathbf{v}^-.$$

Then :

$$\nabla \mathbb{1}^\pm = \pm \mathbf{n} \delta, \quad \nabla \delta = \mathbf{n} \delta'.$$

and then

$$\nabla \mathbf{v} = (\nabla \mathbf{v})^- \mathbb{1}^- + (\nabla \mathbf{v})^+ \mathbb{1}^+ + [[\mathbf{v}]] \otimes \mathbf{n} \delta.$$

General result, curved interface, discontinuous viscosity :

$$[[\rho \mathbf{v} \cdot \mathbf{n}]] = 0 \quad [[\sigma \mathbf{n}]] = \frac{2\gamma}{R} \mathbf{n} - \nabla_T \gamma$$

with :

$$\gamma = -2 [[\nu]] \rho \mathbf{v} \cdot \mathbf{n} \quad \text{surface tension}$$

$$\frac{2}{R} = \nabla_T \cdot \mathbf{n} = \text{total surface curvature}$$

$$\nu = \int \eta \partial_z (1/\rho) \, dz \quad \text{the surface intrinsic property}$$

On a plane interface with a continuous viscosity :

$$\begin{aligned} [[\rho v_z]] &= 0 & [[\sigma_{zz}]] &= 0 \\ [[v_x]] &= 0 & [[\sigma_{xz}]] &= \partial_x (2\eta [[v_z]]) \end{aligned}$$

Numerical test

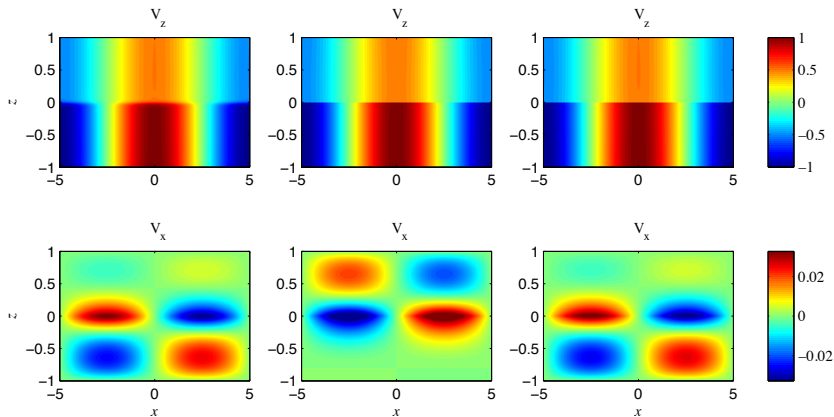
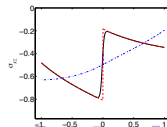


Figure: Velocities for the solutions as function of x and z . Left panels : numerical solution for a diffuse interface. Middle : numerical solution for a sharp interface by using the 'classical' jump conditions (continuous traction). Right : numerical solution for a sharp interface by using the 'classical' jump conditions (continuous traction).



Conclusion : when there is a mass transfert across a density jump interface, then the traction is not continuous. There is a 'dynamic' surface tension.

Prospect

- Inner-core convection
- Mantle convection (410 and 660 km phase changes)
- Rayleigh-Benard and Rayleigh-Taylor convection with phase change
- More generally fluid dynamics with phase change
- Slichter mode ?
- P-SV conversion ?
- (spherical) Shock waves ?
- Energy, temperature, entropy jumps ?
- Mesure of the surface tension ?



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Internal geophysics (Physics of Earth's interior)

Jump conditions and dynamic surface tension at permeable interfaces such as the inner core boundary



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Earth's Inner core special issue