

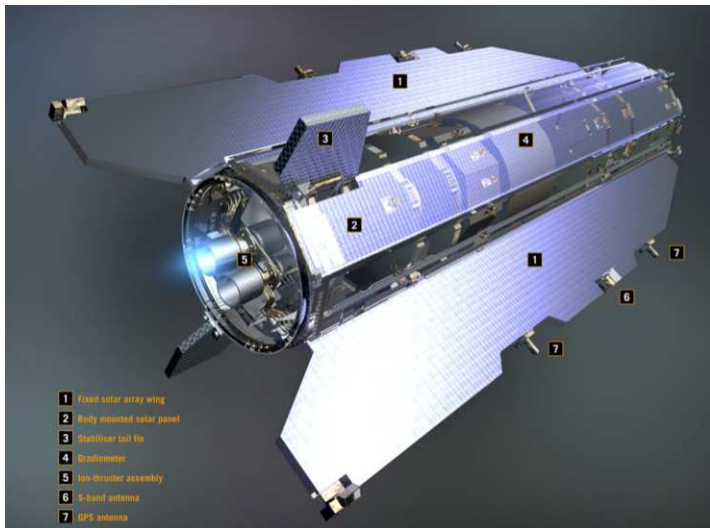
# Comparison of the global gravitational effects obtained respectively from LITHO1.0 geophysical model and GOCE satellite data

Jérôme Verdun<sup>(1)</sup>, Clément Roussel<sup>(1)</sup>, José Cali<sup>(1)</sup>, Frédéric Masson<sup>(2)</sup>

1: GeF Lab - Geomatics and Geosciences Team (L2G) Cnam ESGT - Le Mans, France

2: Institute of Physics of the Earth of Strasbourg (IPGS) - EOST, Strasbourg, France



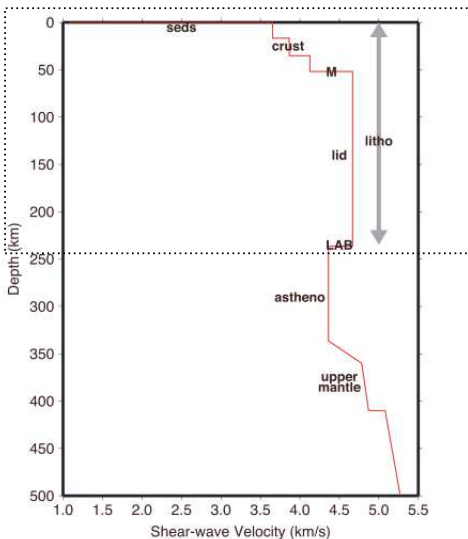


GOCE Level 2 Product Data Handbook, EGG-C, 2014

GOCE  $\Rightarrow$  full tensor gravity gradients = direct, global, high spatial resolution ( $\sim 100$  km), 1 mE accuracy

# LITHO1.0

Available at 1°x1° resolution



Pasyanos, M. E., T. G. Masters, G. Laske, and Z. Ma (2014), LITHO1.0: An updated crust and lithospheric model of the Earth, *J. Geophys. Res. Solid Earth*, 119, 2153–2173, doi:10.1002/2013JB010626.

# LITHO1.0: physical parameters

## Starting models

num	layer	parameters	
1	ice	(thickness), (vp), (vs), ( $\rho$ ), (Q)	Surface
2	water	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
3	sediment layer 1	(thickness), (vp), (vs), ( $\rho$ ), (Q)	Basement
4	sediment layer 2	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
5	sediment layer 3	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
6	upper crust	thickness, vp, vs, $\rho$ , (Q)	Moho
7	middle crust	thickness, vp, vs, $\rho$ , (Q)	
8	lower crust	thickness, vp, vs, $\rho$ , (Q)	
9	lithospheric mantle (lid)	thickness, vp, vs, ( $\rho$ ), (Q)	LAB
10	asthenospheric mantle	vp, vs, ( $\rho$ ), (Q)	

ak135

## CRUST1.0

Laske et al., 2012

## LLNL-G3D

Simmons et al., 2012

Pasyanos, 2010 (thickness)

Kennett et al., 1995

<sup>a</sup>Parameters unmodified from the starting model are shown in parentheses.

Pasyanos et al., 2014

# LITHO1.0: physical parameters

## Starting models

num	layer	parameters	
1	ice	(thickness), (vp), (vs), ( $\rho$ ), (Q)	Surface
2	water	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
3	sediment layer 1	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
4	sediment layer 2	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
5	sediment layer 3	(thickness), (vp), (vs), ( $\rho$ ), (Q)	
6	upper crust	thickness, vp, vs, $\rho$ , (Q)	Basement
7	middle crust	thickness, vp, vs, $\rho$ , (Q)	
8	lower crust	thickness, vp, vs, $\rho$ , (Q)	
9	lithospheric mantle (lid)	thickness, vp, vs, ( $\rho$ ), (Q)	Moho LAB
10	asthenospheric mantle	vp, vs, ( $\rho$ ), (Q)	

ak135

<sup>a</sup>Parameters unmodified from the starting model are shown in parentheses.

Pasyanos et al., 2014

## CRUST1.0

Laske et al., 2012

## LLNL-G3D

Simmons et al., 2012

Pasyanos, 2010 (thickness)

Kennett et al., 1995

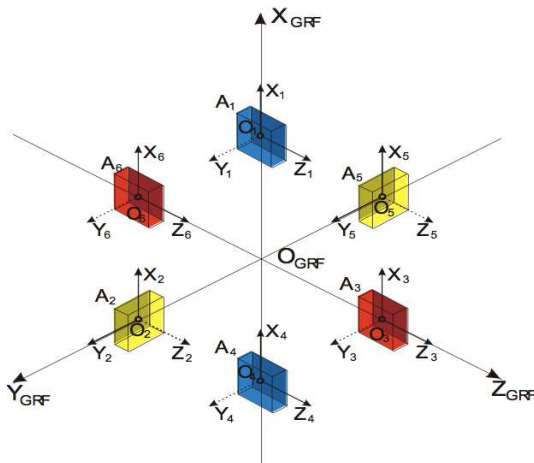
Possible comparison: FTG components deduced respectively from forward modelling and GOCE data

# Outline

- 1 Numerical computation of Earth gravity gradients
- 2 Overview of first results
- 3 Work in progress and future work

# Brief review of GOCE reference frames

## Gradiometer Reference Frame (GRF)

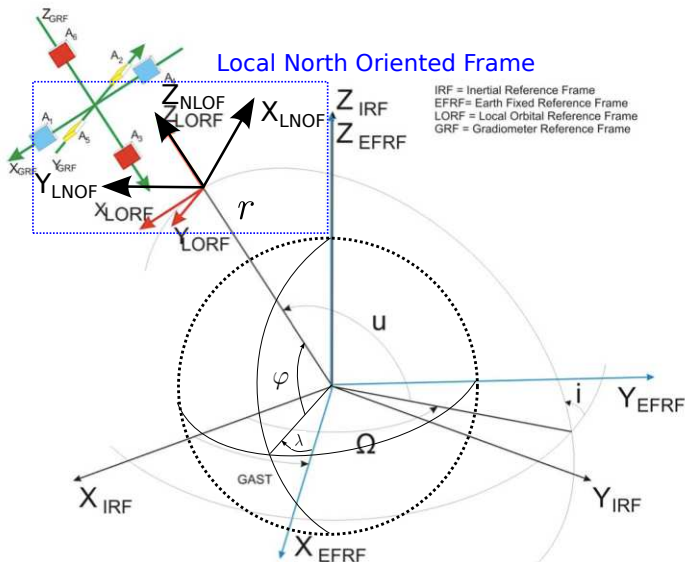


GOCE Level 2 Product Data Handbook, EGG-C, 2014

## Accelerometer Reference Frames (ARF)

.....► less sensitive axis

# Brief review of GOCE reference frames



Original figure from: GOCE Level 2 Product Data Handbook, EGG-C, 2014

# Computation of gravity gradients in the LNOF

# Computation of gravity gradients in the LNOF

Possible by means of numerical integration given a geophysical model

# Computation of gravity gradients in the LNOF

Possible by means of numerical integration given a geophysical model

- 1 **Tesseroid**<sup>1</sup>: Gauss-Legendre Quadrature based on spherical prisms ;

# Computation of gravity gradients in the LNOF

Possible by means of numerical integration given a geophysical model

- ① **Tesseroid**<sup>1</sup>: Gauss-Legendre Quadrature based on spherical prisms ;
- ② **Our own software**<sup>2</sup>: Gauss-Legendre Quadrature based on ellipsoidal prisms.

# Computation of gravity gradients in the LNOF

Possible by means of numerical integration given a geophysical model

- 1 **Tesseroid**<sup>1</sup>: Gauss-Legendre Quadrature based on spherical prisms ;
- 2 **Our own software**<sup>2</sup>: Gauss-Legendre Quadrature based on ellipsoidal prisms.

- 1: Uieda, L., *Tesseroids: forward modeling of gravitational fields in spherical coordinates*: figshare, 2014.
- 2: Roussel, C., Verdun, J., Cali, J. and F. Masson, *Complete gravity field of an ellipsoidal prism in spherical coordinates by Gauss-Legendre quadrature*: in prep for submission to Geophysical J. Int., 2014.

# Basic notations for the gravity field

# Basic notations for the gravity field

① Gravitational potential: 
$$V = \iiint_{Terre} \frac{G \rho(M) d\tau_M}{MP},$$
$$G = 6,67 \times 10^{-11} \text{ m}^3.\text{kg}^{-1}.\text{s}^{-2}$$

# Basic notations for the gravity field

① Gravitational potential:  $V = \iiint_{Terre} \frac{G \rho(M) d\tau_M}{MP},$   
 $G = 6,67 \times 10^{-11} \text{ m}^3.\text{kg}^{-1}.\text{s}^{-2}$

② Gravitational acceleration:

$$\mathbf{g} = \nabla_P (V) = \begin{bmatrix} V_x \\ V_y \\ V_z \end{bmatrix}, \quad V_\alpha = \partial_\alpha V, \quad \alpha = x, y, z$$

# Basic notations for the gravity field

① Gravitational potential:  $V = \iiint_{Terre} \frac{G \rho(M) d\tau_M}{MP},$   
 $G = 6,67 \times 10^{-11} \text{ m}^3.\text{kg}^{-1}.\text{s}^{-2}$

② Gravitational acceleration:

$$\mathbf{g} = \nabla_P (V) = \begin{bmatrix} V_x \\ V_y \\ V_z \end{bmatrix}, \quad V_\alpha = \partial_\alpha V, \quad \alpha = x, y, z$$

③ Gravity gradient tensor (Marussi's tensor)

$$\mathbf{T} = \nabla_P (\nabla_P (V)) = \begin{bmatrix} V_{xx} & V_{xy} & V_{xz} \\ V_{yx} & V_{yy} & V_{yz} \\ V_{zx} & V_{zy} & V_{zz} \end{bmatrix}, \quad V_{\alpha\beta} = \partial_\beta V_\alpha = \partial_\beta (\partial_\alpha V), \quad \alpha, \beta = x, y, z$$

# Basic notations for the gravity field

① Gravitational potential:  $V = \iiint_{Terre} \frac{G \rho(M) d\tau_M}{MP}$ ,  
 $G = 6,67 \times 10^{-11} \text{ m}^3.\text{kg}^{-1}.\text{s}^{-2}$

② Gravitational acceleration:

$$\mathbf{g} = \nabla_P (V) = \begin{bmatrix} V_x \\ V_y \\ V_z \end{bmatrix}, \quad V_\alpha = \partial_\alpha V, \quad \alpha = x, y, z$$

③ Gravity gradient tensor (Marussi's tensor)

$$\mathbf{T} = \nabla_P (\nabla_P (V)) = \begin{bmatrix} V_{xx} & V_{xy} & V_{xz} \\ V_{yx} & V_{yy} & V_{yz} \\ V_{zx} & V_{zy} & V_{zz} \end{bmatrix}, \quad V_{\alpha\beta} = \partial_\beta V_\alpha = \partial_\beta (\partial_\alpha V), \quad \alpha, \beta = x, y, z$$

Symmetric tensor with null trace out of masses.

# Spherical vs ellipsoidal prism

# Spherical vs ellipsoidal prism

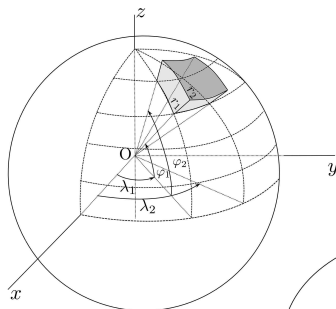
- Two mass elements for decomposing the Earth body

# Spherical vs ellipsoidal prism

- Two mass elements for decomposing the Earth body
- Ellipsoidal prisms are well-suited for modelling the flattening of the Earth

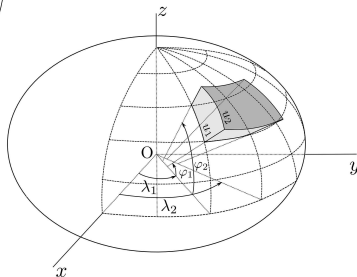
# Spherical vs ellipsoidal prism

- Two mass elements for decomposing the Earth body
- Ellipsoidal prisms are well-suited for modelling the flattening of the Earth



Spherical prism

Ellipsoidal prism



# Gauss-Legendre Quadrature

# Gauss-Legendre Quadrature

Prism domain

$$\int_{\lambda_1}^{\lambda_2} \int_{\varphi_1}^{\varphi_2} \int_{u_1}^{u_2} f(\lambda_P, \theta_P, r_P, \lambda_S, \varphi_S, u_S) du_S d\varphi_S d\lambda_S$$

Any gravity related quantity

Observation point (spherical coordinates)

Source point (ellipsoidal coordinates)

# Gauss-Legendre Quadrature

Prism domain

Any gravity related quantity

$$\int_{\lambda_1}^{\lambda_2} \int_{\varphi_1}^{\varphi_2} \int_{u_1}^{u_2} f(\lambda_P, \theta_P, r_P, \lambda_S, \varphi_S, u_S) du_S d\varphi_S d\lambda_S$$

Observation point (spherical coordinates)

Source point (ellipsoidal coordinates)

$\approx$

# Gauss-Legendre Quadrature

Prism domain

$$\int_{\lambda_1}^{\lambda_2} \int_{\varphi_1}^{\varphi_2} \int_{u_1}^{u_2} f(\lambda_P, \theta_P, r_P, \lambda_S, \varphi_S, u_S) du_S d\varphi_S d\lambda_S$$

Any gravity related quantity

Observation point (spherical coordinates)

Source point (ellipsoidal coordinates)

$\approx$

Prism domain

$$\frac{(\lambda_2 - \lambda_1)(\varphi_2 - \varphi_1)(u_2 - u_1)}{8} \sum_{i=1}^{n_\lambda} \sum_{j=1}^{n_\varphi} \sum_{k=1}^{n_u} \omega_i \omega_j \omega_k f(\lambda_P, \theta_P, r_P, \lambda_{S_i}, \varphi_{S_j}, u_{S_k})$$

Quantity to be integrated

Observation point (spherical coordinates)

Source point at Gauss-Legendre nodes (ellipsoidal coordinates)

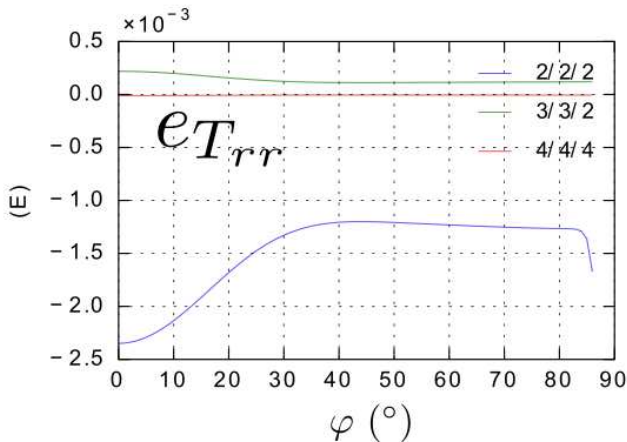
Gauss-Legendre weighing coefficients

# Gauss-Quadrature validation: example for $V_{zz}$

$e_{T_{rr}} = e_{V_{zz}} = (V_{zz})_{GLQ} - (V_{zz})_{exact formulas}$  for an ellipsoidal layer of constant density.

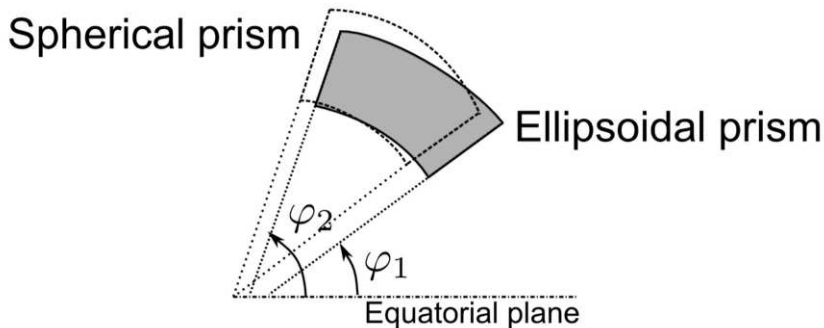
# Gauss-Quadrature validation: example for $V_{zz}$

$e_{T_{rr}} = e_{V_{zz}} = (V_{zz})_{GLQ} - (V_{zz})_{exact\ formulas}$  for an ellipsoidal layer of constant density.



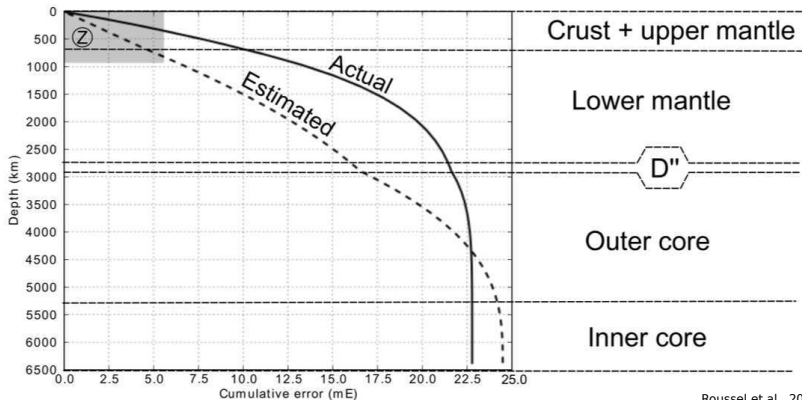
Roussel et al., 2014

# To what extent the spherical approximation remains valid ?

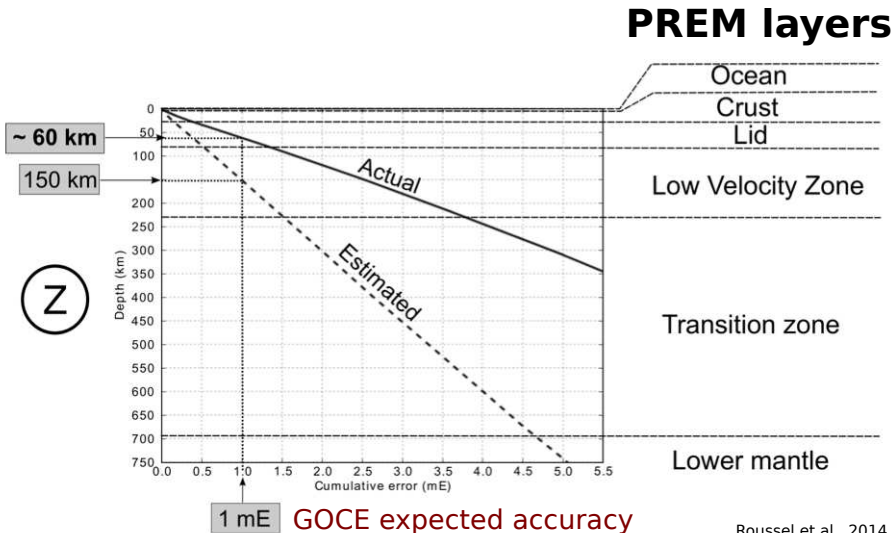


# To what extent the spherical approximation remains valid ?

## PREM layers



# To what extent the spherical approximation remains valid ?

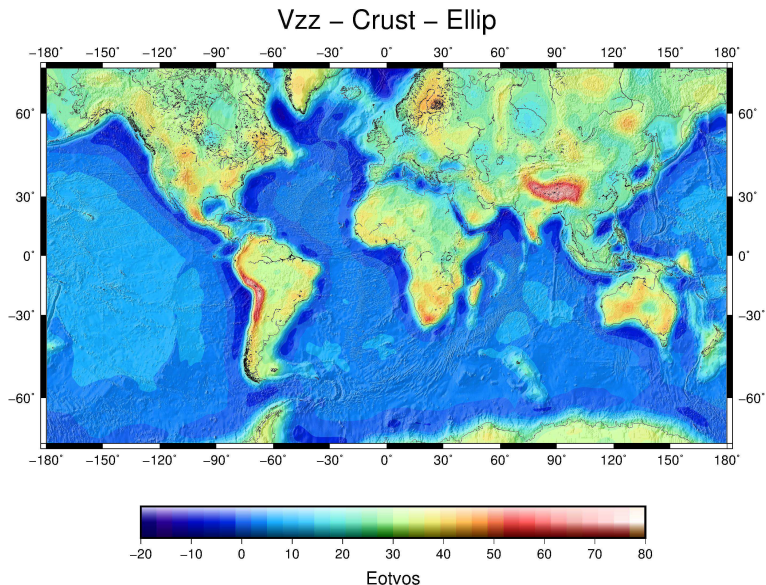


# To what extent the spherical approximation remains valid ?

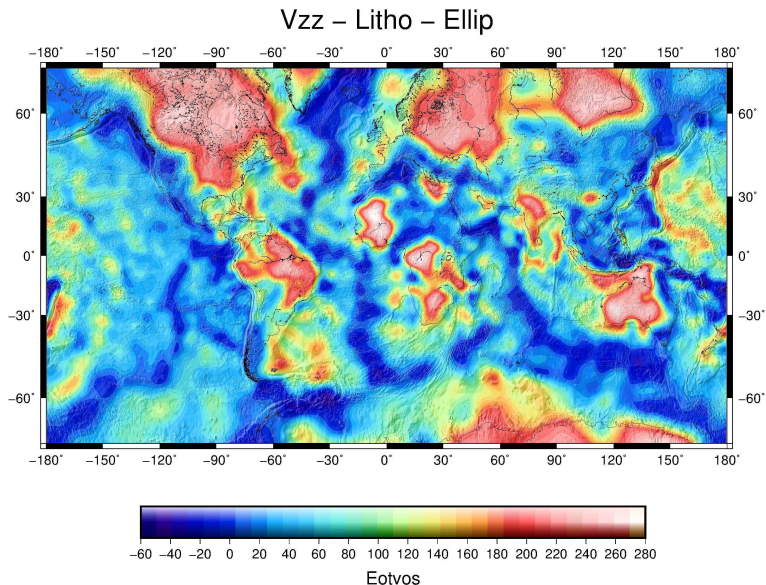
## Conclusions

The numerical computation of global gravity gradient tensor up to the lithospheric mantle suitable with GOCE accuracy has to be performed with ellipsoidal prisms.

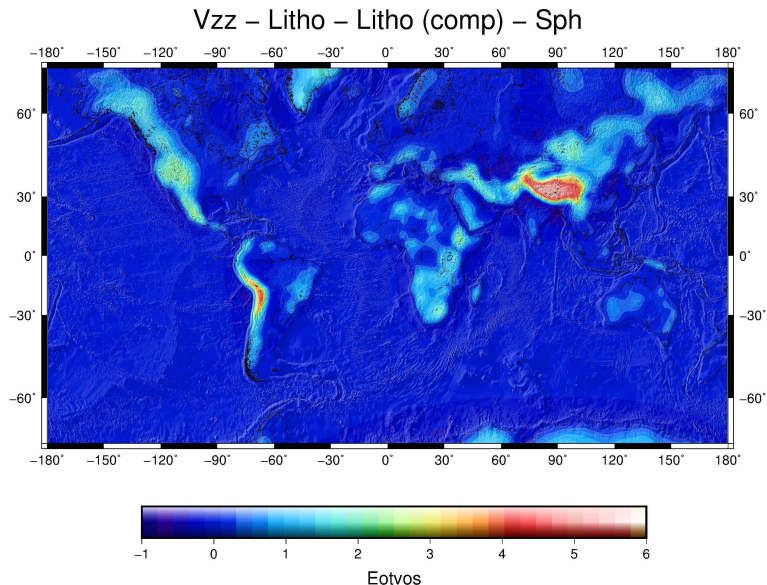
# CRUST1.0 gravitational effect



# LITHO1.0 gravitational effect

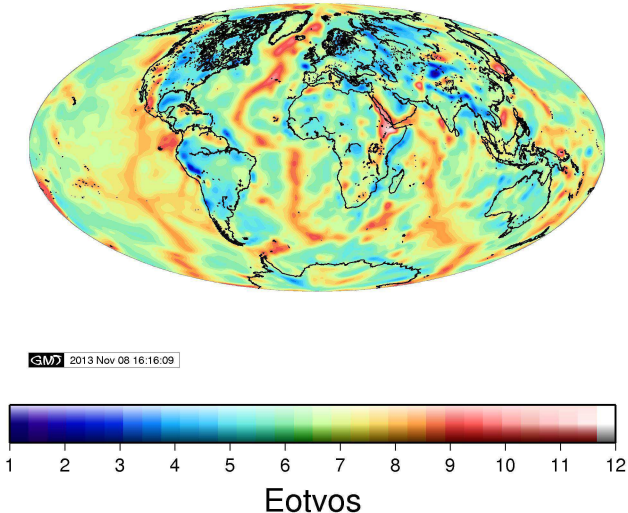


# LITHO1.0 - LITHO1.0 after Airy isostatic compensation



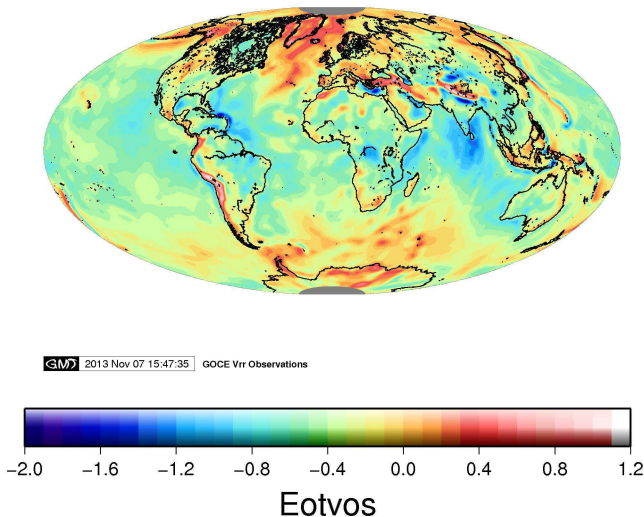
# Anomaly of CRUST1.0/PREM combined model with respect to PREM

Anomaly of combined CRUST1.0-PREM model/PREM



# Anomaly of GOCE 1 year $V_{zz}$ data with respect to PREM

Anomaly of 1 year  $V_{zz}$  GOCE data /PREM



# Conclusions and prospects

- 1 The need to use ellipsoidal prisms in the computation of global gravity effects produced by the whole lithosphere at 1 mE accuracy has been demonstrated conclusively.

# Conclusions and prospects

- 1 The need to use ellipsoidal prisms in the computation of global gravity effects produced by the whole lithosphere at 1 mE accuracy has been demonstrated conclusively.
- 2 The software using ellipsoidal prism to perform GLQ integration has been validated and is now efficient.

# Conclusions and prospects

- 1 In our experiment, the direct computation of gravity gradients by interpolation of available GOCE data leads to underestimated gravity anomaly values; further investigation must be carried out.

# Conclusions and prospects

- ① In our experiment, the direct computation of gravity gradients by interpolation of available GOCE data leads to underestimated gravity anomaly values; further investigation must be carried out.
- ② Thanks to GOCE spectral sensitivity, suitable calculation for the purpose of comparison with geophysical models must necessarily include the lithosphere.

# Conclusions and prospects

- 1 Computation of the gravitational effect produced by LITHO1.0-PREM combined model.

# Conclusions and prospects

- 1 Computation of the gravitational effect produced by LITHO1.0-PREM combined model.
- 2 Computation of GOCE gravity gradient anomalies using 1 year GOCE data recently available as evenly sampled grids (GRD SPW 2)

Thank you for your kind attention!

